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Pre-Columbian jadeitite artifacts from the Golden Rock Site, St. Eustatius, Lesser Antilles, with special reference to jadeitite artifacts from Elliot's, Antigua: Implications for potential source regions and long-distance exchange networks in the Greater Caribbean



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1 **Highlights**

- 2 ➤ A detailed electron microscopy (microprobe) analysis and imaging of pre-Columbian
3 jadeitite celts, St Eustatius, Lesser Antilles, is offered for the first time.
- 4 ➤ The mineralogical and petrological data are compared with data from samples of the known
5 source regions of Caribbean jadeitite in Guatemala, Cuba and Dominican Republic.
- 6 ➤ The provenance analysis suggests a Guatemalan source, strengthening the view of long-
7 distance exchange networks in the Caribbean since the Early Ceramic Age.

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3 **artifacts from Elliot's, Antigua: Implications for potential source**
4 **regions and long-distance exchange networks in the Greater**
5 **Caribbean**

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26 **Keywords**

27 Jadeitite, Jade, pre-Columbian artifacts, exchange networks, Caribbean.

28 **Abstract**

29 A detailed electron microscopy analysis of jadeitite celts from the Early Ceramic Age Golden
30 Rock settlement on the small volcanic island of St Eustatius, Lesser Antilles, is presented in an
31 effort to identify the source region(s) of these jadeitite axes and evaluate the extent of trade
32 networks in the Caribbean during pre-Columbian times through which those tools (or source rocks)
33 circulated. Habitation at the site occurred between ca. AD 230 - 890, and the jadeitite tools most
34 likely date between cal. AD 600 and 825/890. We argue that in provenancing jadeitite emphasis
35 should be placed on the identification of the entire mineral assemblage (including the accessory
36 minerals) and textures, given the complex geological histories and processes that form this quasi-
37 monomineralic rock. Indeed, the mineral assemblages and the characteristics of the individual
38 minerals within the studied jadeitite samples are far from homogeneous, suggesting either the
39 source has a high degree of internal variation or there are multiple sources. We have identified three
40 jadeitite groups among the analyzed samples on the basis of mineralogical assemblages: Group 1
41 consists of samples bearing phlogopite (plus phengite and epidote but without lawsonite or quartz);
42 Group 2 contains lawsonite (plus phengite and quartz but without phlogopite and epidote); and
43 Group 3 is defined by containing hematite. Importantly, we have not identified glaucophane in the
44 analyzed materials. The comparison of these artifact data with those of jadeitites from the four
45 potential known source regions of Caribbean jade, i.e., Guatemala (North and South Motagua Fault
46 mélanges), Cuba (Sierra del Convento mélange) and Dominican Republic (Río San Juan complex)
47 allows us to conclude that both the North and South Motagua Valley mélanges are the most likely
48 sources for artifact groups 2 and 3, and perhaps also for group 1. This identification supports earlier
49 hypotheses on the existence of pan-regional exchange networks, tying the northern Antilles with the

50 Isthmo-Colombian region (Colombia, Panama and Costa Rica) during the Ceramic Age of the
51 Caribbean.

52 **1 Introduction**

53 Caribbean archaeology has recently experienced an increased interest in the study of inter-
54 island and pan-regional interactions (e.g., Berman, 2011; Hofman and Hoogland, 2011; Hofman et
55 al., 2007, 2011; Rodríguez Ramos and Pagán Jiménez, 2006). One of the major lines of research has
56 focused on the study of the exchange of exotics between the islands and the existence of regional
57 and macro-regional exchange networks (Berman, 2011; Cody, 1991, 1993; Hardy, 2008; Hofman et
58 al., 2007; Knippenberg, 2004, 2006, 2007; Rodríguez Ramos, 2002, 2010, 2011). Research on lithic
59 artifacts plays a key role in this endeavor, as it is the material category that can provide very
60 detailed insights into the entire systemic process of acquiring, making, distributing, and using tools
61 and objects (Knippenberg, 2006; Torrence, 1986). Among the transacted materials in these
62 networks is jadeitite, a rock primarily composed of the sodic clinopyroxene jadeite. Jadeitite is well-
63 known due to the very important role it fulfilled and broad usage it had in the great Meso- and
64 Central American civilizations, including the Olmec, Maya, Aztec and Costa Rican cultures (Jones,
65 1998; Lange, 1993). The more limited use of jadeitite in the Caribbean had previously been
66 hypothesized from rare finds of green lithic artifacts resembling this semi-precious stone (Chanlaitte
67 Baik, 1984; Chanlaitte Baik and Narganes Storde, 1990; Fewkes, 1922; Haag, 1970; Sued Badillo,
68 1979). However, more recent work has shown that most of the personal adornments were
69 erroneously identified, as these are mainly composed of nephrite and serpentinite (Cody, 1991;
70 Murphy et al., 2000; Watters and Scaglione, 1994). Initial X-ray diffraction studies conducted on a
71 Bahamian celt by Rose (1987) together with the in-depth petrographic and mineralogical work by
72 Harlow et al. (2006) on a number of celts (axes) and celt fragments from the Elliot's site on the
73 island of Antigua identified with high resolution techniques the use of jadeitite for the first time in
74 the Antilles.

75 The identification of jadeitite artifacts in the Caribbean has had profound implications for our
76 understanding of Antillean interaction networks operating in pre-Columbian times because, until
77 recently, the closest jadeitite sources were only known in the Motagua Valley in Guatemala
78 (Harlow, 1993, 1994; Harlow et al., 2006). Close similarity between the Antiguan artifacts and
79 Guatemalan source material allowed both Rose (1987) and Harlow et al. (2006) to propose this
80 distant Meso-American source region as being of significance to the pre-Columbian cultures of the
81 Caribbean. Subsequent work by Harlow on greenstone celts and adzes from Tecla I and II, Punta
82 Candelero, and Río Tanamá on Puerto Rico (Rodríguez Ramos, 2007) and Judith Fancy and Aklis

83 in the Virgin Islands (Hardy, 2008) confirmed the common presence of jadeitite in the
84 archaeological record of the Antilles and its close association with the use as raw material for axes
85 and much rarer adzes. Combined with iconographic, botanical, metallurgical, technological and
86 stylistic concomitances between the Isthmo-Colombian area – particularly Costa Rica, Panama, and
87 Colombia – and Early Ceramic Age northeastern Caribbean cultures, initial ideas on a distribution
88 trajectory from Guatemala via the Isthmo-Colombian area directly towards the northeastern Antilles
89 (Puerto Rico) were formulated (Rodríguez Ramos, 2007, 2010, 2011).

90 Harlow et al. (2006), however, noted some mineralogical/petrological dissimilarity between
91 the Antiguan artifacts and the South Motagua *mélange*, and indicated that "*yet unknown sources in*
92 *a similar geological terrane (serpentinite–matrix *mélange*) on Cuba, Hispaniola or Jamaica cannot*
93 *be ruled out*". In fact, since that study was published, two new sources of jadeitite have been
94 discovered in the Greater Antilles (Figure 1), in the Río San Juan region, northern Cordillera,
95 Dominican Republic (Schertl et al., 2007, 2012) and in the Sierra del Convento, Guantanamo
96 province, SE Cuba (Cárdenas-Párraga et al., 2010, 2012; García-Casco et al., 2009). These recent
97 discoveries are extremely important for archaeological studies in the region. First, because they
98 provide potential local sources for Caribbean jadeitite artifacts, including perhaps those that have
99 been located in archaeological contexts as far as the Lesser Antilles. Second, because they represent
100 only two out of a larger number of potential geological deposits in the region, suggesting that the
101 discovery of new geologic deposits cannot be ruled out (e.g., in central Cuba, Jamaica, northern
102 Venezuela and Colombia; Figure 1). The possibility of finding new sources in the region makes the
103 task of firmly establishing the provenance of jadeitite artifacts in the Greater Caribbean very
104 challenging. However, not positively identifying a source should not detract from the effort, as a
105 negative result may suggest an additional unknown source, which by itself would be of major
106 significance for the archaeology of the region.

107 One of the main current issues regarding the use of jadeitite, therefore, centres on the roles
108 that these newly discovered Antillean sources played in the pan-regional networks where artifacts
109 made of this raw material circulated in the Greater Caribbean. In order to address the
110 aforementioned issues, this work presents the results of an in-depth mineralogical, petrological, and
111 geo-chemical analysis of jadeitite artifacts from the site of Golden Rock in St. Eustatius to
112 determine their geological provenance by comparing their signatures with those documented in the
113 different potential source regions.

114 **1.1 Some facts about jade**

115 The gemological term jade collectively encompasses two different types of semiprecious
116 rocks that are different mineralogically and in geologic origin. Jadeite jade (or jadeitite) is a quasi-

117 monomineralic rock made of jadeite, a sodic clinopyroxene, while nephrite jade (or
118 actinolite/tremolite), the classical Chinese jade, is composed essentially of actinolite or tremolite,
119 both calcic clinoamphiboles. The latter is relatively common in nature, whereas jadeitite is rare.
120 Only ca. 20 localities for jadeitite are known worldwide (e.g., Tsujimori and Harlow, 2012),
121 including the four located in the Caribbean realm, namely the North Motagua Fault region of
122 Guatemala, the South Motagua Fault region of Guatemala, the Rio San Juan of Dominican
123 Republic, and the Sierra del Convento of eastern Cuba. There are two additional localities in North
124 America: New Idria in California and the Vizcaino peninsula in Baja California.

125 All jadeitite deposits are related to geological complexes that have undergone high pressure
126 due to subduction of lithospheric plates, and most are hosted in ultramafic rock complexes formed
127 in the Earth's mantle at > 30 km depth. Fluids released by subducting oceanic lithosphere at great
128 depth interact with upper plate ultramafic mantle rocks (i.e., mantle wedge) forming serpentinites
129 and, locally, jadeitite deposits by hydrothermal deposition in veins and/or metasomatic
130 transformation of other adjacent rock types (Harlow and Sorensen, 2007; Tsujimori and Harlow,
131 2012). These ultramafic complexes are characterized by tectonic mélanges consisting of serpentinite
132 (essentially, antigorite) matrix and tectonic blocks of high-pressure rocks such as blueschist,
133 eclogite, garnet-epidote amphibolite and jadeitite, among others. This high-pressure metamorphic
134 characteristic is important for understanding archaeological jadeitite distributions in the Caribbean
135 (and elsewhere) because, although ultramafic rock complexes are abundant in the region (e.g.,
136 Lewis et al., 2006), only some of them experienced high-pressure in subduction zones and thus are
137 potential sites for hosting jadeitite deposits (Figure 1).

138 **2 Archaeological context**

139 The jadeitite artifacts discussed in this paper were excavated at Golden Rock, a large Ceramic
140 Age settlement site centrally located on the small volcanic island of St. Eustatius, one of
141 Netherlands' overseas municipalities in the Caribbean (Figure 2). Following earlier work in the
142 1920s by de Josselin de Jong (1947), who attested to the significance of the site, rescue excavation
143 work in the 1980s by Versteeg and Schinkel (1992) from Leiden University unearthed the remains
144 of a multi-phase Amerindian settlement. The findings include six different house plans flanked by
145 several smaller structures and an associated circumscribed refuse midden deposit, all within an area
146 of approximately 1200 m². The investigated area only represents a portion of the original settlement
147 (Figure 2; Versteeg and Schinkel, 1992).

148 The systematically excavated midden deposit, ca. 400m² in size, produced a wealth of
149 artifacts made from pottery, lithics, shell, and coral, as well as enormous quantities of food debris,

150 comprising predominantly shells, small fish and rodent bones. Multiple radiocarbon dates showed
151 that the site was inhabited between cal. AD 230 - 890 making it the oldest Ceramic Age site on this
152 small volcanic island. Dated hearths (h1-h3) situated underneath the artifact rich midden deposit
153 suggest refuse material started to accumulate there somewhere between cal AD 230 – 660 (Versteeg
154 and Schinkel, 1992, 204). On the basis of dates from the associated house structures, its most
155 significant period of use, however, is later and must have been between AD 600 and 825-890,
156 (Versteeg and Schinkel, 1992, 229; see also Knippenberg, 2006). This is likely the period when the
157 jadeitite artifacts of the site were used and this usage is probably for a large part contemporaneous
158 with that at Elliot's on Antigua. It should be noted that although in their original paper Harlow et al
159 (2006) did not provide specific dates for the use of the jadeitite on Antigua, as the Elliot's site had
160 not been absolutely dated, the description of the ceramic assemblage associated to these jadeitite
161 celts provided in Murphy's (2001, 84-174) dissertation indicates that these are associated to a
162 younger cultural component than that documented in the neighboring radiocarbon dated Royall's
163 site (cal AD 250 – 630), therefore most likely dating to the later part of the Early Ceramic Age (AD
164 400 – 800).

165 Of the 3238 lithic artifacts found in the refuse and comprising the entire stone tool inventory
166 and related manufacturing debris associated with the different Amerindian households identified at
167 the site, seven axe fragments were classified as possibly being made from jadeitite, making up ca.
168 8% of all axes recovered. The jadeitite pieces include three edge flakes, another flake with a ground
169 dorsal face, as well as two medial-lateral and one proximal-lateral axe fragment (Table 1). All
170 jadeitite artifacts came from the top two thirds of an on average 60 to 70 cm thick midden deposit.
171 Recorded depths of the specimen varied between 0 and 40 cm below the surface. Within the midden
172 deposit no distinction could be made between different occupation phases; the finding of the
173 jadeitite in the top 40 cm suggests that most likely this material should not be associated with the
174 earliest use of the midden area. All of the identified jadeitite pieces came from finished tools,
175 indicating that local manufacture of jadeitite axes at Golden Rock had not occurred. Furthermore,
176 the items still possessing (part of) the edge exhibit macroscopic use-wear types that are typically
177 found on axes, suggesting these jadeitite tools had likely been used as wood chopping implements.
178 From these seven specimens, four were selected for in-depth petrographic, mineralogical, and geo-
179 chemical analysis (see Table 2 and figure 2).

180 Apart from these jadeitite axe fragments, the Golden Rock site produced a wealth of other
181 axes and axe manufacture related debris made from a range of rock materials. Predominant is a
182 fine-grained grey-green tuffaceous mudstone originating from St. Martin, making up almost 73% of
183 the total axe inventory. High amounts of manufacture debris in the form of flakes, shatter, and pre-

184 forms, as well as fragments of finished pieces (ca. 20% of all found lithics) shows that Golden Rock
185 was involved in the on-site production of celts of this raw material (Knippenberg, 2006). In addition
186 the inventory includes rare fragments of finished axes made from fine-grained mafic igneous rocks
187 and different metamorphic rocks other than jadeitite.

188 **3 Analytical and computation techniques**

189 The provenancing of jadeitite has proven to be a very challenging task (see Lange, 1993 and
190 contributions therein). Unlike other lithic sourcing studies (e.g., obsidian, chert) where the use of
191 geochemical techniques for determining trace-element compositions and fingerprinting artifacts has
192 proven highly effective, jadeitite appears to be less suitable for such approach (Harlow, 1993). As a
193 consequence of the common occurrence of recurrent metasomatic (i.e., chemical infiltration and
194 reaction) and vein filling-jadeite forming events in a specific rock mass involving fluids (i.e.,
195 chemicals) of varied sources and fluid paths, jadeitite rocks are not likely to cluster into narrow or
196 highly restricted zones in plots of multidimensional chemical space.

197 This notion has guided the methodology of the present work and that of the earlier Antigua
198 study (Harlow et al., 2006). The primary objective of the current work therefore is to determine the
199 entire mineralogical constitution of the jadeitite artifacts and to characterize the important mineral
200 phases. This has been done by a combination of analytical and computation techniques, including
201 optical (polarized light) microscopy and electron beam examination (microprobe; WDS, EDS, and
202 X-ray imaging) as well as projection and condensation of the multidimensional chemical space in
203 order to produce tetrahedral diagrams (ACFN deluxe diagram) containing the full mineralogical
204 information of the samples (Appendix A). Representative chemical analyses and structural formulas
205 of the phases are given in table A1

206 The four samples that were analyzed (Figure 3) were cut to prepare polished thin sections for
207 microscopy and electron-probe analysis. As described below, all samples are jadeite jade (jadeitite).
208 Both, sodic clinopyroxene, jadeite, and sodic-calcic clinopyroxene, omphacite, are present in all
209 samples except GR4-FIII.4 which contains only jadeite. The detailed mineralogical descriptions
210 given below are important for both archaeological and geological reasons, as they can be used as a
211 reference for future descriptions of Caribbean jadeitite artifacts and geological source samples,
212 including any new sources in the region.

213 4 Textures and mineral composition

214 4.1 Sample GR24-B-III.1

215 Sample GR24-B-III.1 (Figure 3A) shows variations in green color as a result of variation in
216 the concentration in jadeite, omphacite and, to a lesser extent, white mica and other minerals.
217 Discrete grains or clots of grains of white mica (phengite) are distinctive visually, particularly with
218 a hand lens. The entire mineral assemblage identified in this sample consists of (in order of
219 abundance; Table 2) jadeite, omphacite, phengite, epidote, titanite and biotite, in addition to minor
220 amounts of apatite, albite and analcime. This sample has a massive heterogranular (fine to coarse
221 grained) texture, in which distinct coarse grained jadeite, clots of omphacite and distinct single
222 crystals or clots of phengite (up to 3 mm in length) are set in a fine-to-medium grained (ca. 0.1-0.5
223 mm in size) matrix of jadeite and, to a lesser extent, omphacite (Figure 4). The textural relationships
224 suggest the following crystallization scheme: $Omp1 \gg Jd \gg Omp2+Phl+Ph+Ep+Ttn \gg Anl+Ab$
225 (where the symbol \gg points along the temporal sequence). The analyzed minerals of this sample
226 are plotted in the ACFN deluxe diagram of Figure 5a.

227 **Jadeite** constitutes much of the rock. The crystals are fine to coarse grained. Backscattered
228 electron (BSE) and X-ray (XR) images show variable composition (Figures 4b, d, e and f and 6a, b
229 and c). The largest grains show an ill-defined concentric zoning, with cores generally richer in Ca,
230 Mg and Fe and rims (and regions close to clots of omphacite and phengite) richer in Na and Al.
231 However, the cores and rims also show patchy zoning. In the cores, Na-Al richer regions are located
232 adjacent to inclusions of corroded omphacite. These regions are, in turn, surrounded by patchy Ca-
233 Mg-Fe richer regions which are also present in the rims, although the outermost jadeite in contact
234 with omphacite and phengite is rich in Na-Al. These relations suggest complex growth of jadeite
235 with recurrent changes in composition from Na-Al rich through Ca-Mg-Fe rich to Na-Al rich
236 (Figure 7). The amounts of calculated of Fe^{3+} and Fe^{2+} are low, indicating jadeite-diopside solid
237 solution with low hedenbergite and aegirine content.

238 The compositions of the pyroxene are presented and classified in the ternary q-jd-aeg diagram
239 (Morimoto et al., 1988; Figure 8). Most of the analyses are classified as jadeite, but some analyses
240 with higher q-component (q >20 mole%) should be described as omphacite following these authors.
241 The phase relations for clinopyroxenes calculated at a reference temperature of 500 °C by Green et
242 al., (2007) have been included in this diagram in order to show the solvus relations in the jadeite-
243 diopside-aegirine system. It should be noted that jadeite (and coexisting omphacite, see below) do
244 not conform to the phase relations calculated at 500 °C in this system (Figure 8), suggesting
245 displacement of the forbidden region due to the effect of temperature of formation (which should be
246 higher than 500 °C) and/or of extra-components (notably FeO, cf. García-Casco et al., 2009) and/or

247 inaccuracy in the model (as suggested by Harlow et al., 2011). However, these point analyses of
248 jadeite with high-q component may also represent cryptic Jd-Omp mixtures not readily discernable
249 in BSE images.

250 **Omphacite** occurs as inclusions within jadeite, as clots dispersed in the matrix and as
251 overgrowths at the edges of jadeite grains (Figures 4 and 6a, b and c). The textural relations indicate
252 that omphacite inclusions represent former single grains that were replaced (corroded) by jd-rich
253 pyroxene, in a first step, and jd-poorer pyroxene in a late stage. However, it cannot be excluded that
254 some of the smaller omphacite grains within jadeite represent late non-lamellar exsolutions
255 (unmixing) rather than inclusions. The clots of omphacite are associated with phengite and, locally,
256 biotite (phlogopite). These clots appear to occupy former "voids" in the rock, as demonstrated by
257 the textural position of the associated phlogopite (see below). However, towards the rims of the
258 clots, omphacite replaces jadeite. Similarly, the fine-grained crystals of omphacite adjacent to
259 jadeite replace the latter and are associated with phengite. The composition of omphacite is diverse
260 but mostly showing jadeite-diopside solid solution with minor contribution of hedenbergite and
261 aegirine components (Figures 7 and 8). The compositions of omphacite with different textural types
262 overlap, but generally the omphacite replacements after jadeite have the most jd-rich compositions
263 while the clots have the most jd-poor compositions. As with jadeite compositions plotting in the
264 "forbidden" jd-omp solvus region, some analyses of omphacite with higher q-component plot in the
265 "forbidden" omphacite-diopside solvus region calculated at 500 °C. This may result from either not
266 considering hedenbergite in the calculated phase relations and/or the temperature of formation being
267 higher than 500 °C.

268 **Phengite** occurs as discrete crystals and clots of crystals dispersed in the rock (Figures 4a and
269 b). As indicated above, it is generally associated with late omphacite (Figures 4c and d), showing
270 an interstitial position denoting a relatively late stage of crystallization. The composition of
271 phengite shows a broad compositional spectrum (Figure 9) with Si up to 6.89 atoms per formula
272 unit indicating solid solution in the muscovite-phengite series with some contributions of paragonite
273 and barian muscovite components. The largest deviations of composition from the muscovite end-
274 member are due to the Tschermak substitution $((\text{Mg,Fe})\text{Si}^{\text{VI}}\text{Al}_1^{\text{IV}}\text{Al}_1)$, implying an important
275 increase in Fe+Mg and Si (Figure 9), as expected for K-white micas crystallized at high-pressure.

276 The presence of Ba-rich K-white mica compositions (up to 3.44 wt% BaO) is a significant
277 feature. The XR map of Ba(L α) performed in this sample shows the chemical zoning at the grain-
278 size scale (Figure 6e). It can be appreciated that the highest concentration in Ba is located towards
279 the borders of the grains, where the contents in Al are also high (Figure 6d), but Si and K decrease
280 (not shown) and Mg and Fe are uncorrelated. This indicates the "celsian-like" substitution

281 $^{[XIII]}Ba^{[IV]}Al^{[XIII]}K_{-1}^{[IV]}Si_{-1}$ instead of the "barian-muscovite" (or "ganterite")
282 substitution $^{[XIII]}Ba^{[VI]}Mg^{[XIII]}K_{-1}^{[VI]}Al_{-1}$, in agreement with Harlow (1995) and Blanco-Quintero et al.
283 (2011). Though not exclusively found in subduction zones, the existence of K-white micas rich in
284 Ba is typical of rocks formed in this environment, interpreted as a result of jadeitite-forming fluids
285 bearing a component derived from the dehydration of subducted oceanic sediments (e.g., Blanco-
286 Quintero et al., 2011; Harlow, 1995). This type of mica has been described in Guatemalan jadeitite
287 from both the North and South Motagua mélanges (Harlow et al., 2011) but not in jadeitites from
288 the Cuban Sierra del Convento or the Dominican Rio San Juan mélanges (Cardenas-Párraga et al.,
289 2012; Schertl et al., 2012).

290 **Phlogopite** is associated with omphacite and with phengite (Figure 6f). The textural position
291 in the center of the clots of omphacite suggests crystallization in voids, reinforcing the view that all
292 the mineral grains from these clots crystallized in cavities filled by hydrothermal fluid. Phlogopite
293 associated with phengite is fine-grained, occupies interstitial positions and appears to be a late-stage
294 alteration of phengite. Similar textural features of phlogopite have been described in jadeitites from
295 the Cuban Sierra del Convento (Cárdenas-Parraga et al., 2012; García-Casco et al., 2009) and the
296 Guatemalan North (common) and South (rarer) Motagua mélanges (Harlow et al., 2011). In these
297 cases, phlogopite is late and appears with or after omphacite, \pm albite, \pm phengite, \pm chlorite in
298 interstitial growths (i.e., late veining but not necessarily cavities). All compositions have similar
299 Mg# (0.80-0.81), but the latter type of phlogopite is poorer in Al, Ba and Na and richer in Si, Fe and
300 Mg than that associated with omphacite (Figures 6f and 10).

301 **Epidote** is present in the clots of omphacite (\pm phengite \pm phlogopite), is fine grained and
302 occupies an interstitial position, suggesting late growth. The composition of epidote has Fe^{3+} in the
303 range 0.47-0.59, which corresponds to 48-61 mole% pistacite component (Figure 5a). **Titanite** is
304 fine to coarse-grained (Figure 4b). The analyzed content of Al is ca. 0.05 apfu. **Plagioclase** appears
305 within jadeite crystals associated to inclusions of omphacite. Its composition is albitic, though with
306 measurable anorthite content (up to 5.5 mole%; Figure 5a). **Analcime** occupies interstices and is a
307 late alteration product of pyroxene; Figure 4c). Its composition is almost stoichiometric
308 $NaAlSi_2O_6 \cdot H_2O$ (Figure 5a).

309 4.2 Sample GR24-J-IV.05

310 Sample GR24-J-IV.05 (Figure 3B) is medium to fine grained, green in color with some
311 variation in the shade of green. The lighter regions reflect the presence of white mica (and albite).
312 The mineral assemblage consists of jadeite, omphacite, epidote, phengite, phlogopite, chlorite,
313 albite, pumpellyite, and titanite (Table 2). This artifact shows an isotropic texture with no preferred
314 orientation of grains, except fine-grained titanite that appears oriented. The texture is heterogranular

315 showing a noteworthy nodular structure (Figures 11a and b). The nodules (up to 0.5 mm in length)
316 consist of omphacite, phengite, jadeite, epidote, albite, \pm phlogopite, \pm chlorite, \pm pumpellyite and are
317 set within a matrix of jadeite and intergrown smaller crystals of jadeite, omphacite and epidote
318 (Figures 11c and d). However, the distribution of crystals in the intergrowths appears to delineate
319 former large single crystals (0.5 mm), probably of omphacite transformed to jadeite+epidote
320 (Figure 11d). The textural relationships suggest the following crystallization sequence: Ttn \gg
321 Omp1 \gg Jd+Ep/Omp2+Phl+Ph+Ep \gg Ab \pm Chl \pm Pmp. Except for the prominent presence of
322 epidote within the intergrowths, the sequence of crystallization, the mineral assemblage, and the
323 textures resemble those of artifact GR24-B-III.1. The analyzed minerals of this sample are
324 presented in the ACFN deluxe diagram of Figure 5b.

325 The composition of **jadeite** is heterogeneous (Figures 7, 8 and 11b and d). The matrix grains
326 enclosing the nodules and the Omp+Ep+Jd intergrowths, and "vein" and "pool" regions that
327 crosscut previous blocky jadeite, show compositions richer in Na and Al (jd component), while
328 jadeite in the intergrowths show the reverse. The low amounts of calculated of Fe³⁺ and Fe²⁺
329 indicate jadeite-diopside solid solution with low hedenbergite and aegirine contents (Figure 8). As
330 with artifact GR24-B-III.1, most of the analyses are jadeite, but some analyses with q component
331 >20 mole% can be described as omphacite. Also, jadeite (and omphacite) with high q-component
332 plot in the "forbidden" solvus region calculated a 500 °C (Green et al., 2007), suggesting formation
333 at higher temperature and/or unaccounted components.

334 The range of **omphacite** compositions is similar to but more restricted than in artifact GR24-
335 B-III.1, probably the result of limited analyses (Figures 7 and 8). However, a higher concentration
336 in Fe²⁺_{total} is apparent (0.15-0.17 apfu, Figure 7) and the amounts of calculated of Fe³⁺ and Fe²⁺
337 indicate jadeite-diopside solid solution with minor contribution of hedenbergite and aegirine
338 components. The grains in the Omp+Ep+Jd replacements are poorer in jd content while the jd-
339 richer compositions are present in the nodules. Most analyses with q-component plot in the
340 "forbidden" solvus region for the omphacite-diopside system at 500 °C with the same explanations
341 as stated earlier.

342 The composition range of **phengite** is wide (Figure 9), indicating solid solution in the
343 muscovite-phengite series (Tschermak substitution) with some contribution of paragonite
344 component. The amount of Fe+Mg and Si is consistent with crystallization at high-pressure. It is
345 noteworthy that Ba-rich compositions have not been detected in this artifact, and that Fe²⁺_{total} is
346 higher than in artifact GR24-B-III.1. Similarly, **phlogopite** has more Fe²⁺_{total} (1.64-1.72 apfu), with
347 lower Mg#, than in artifact GR24-B-III.1, approaching the field of biotite (Figure 10).

348 **Clinozoisite** (Fe^{3+} 0.07.-0.12; Figure 5b) rather than epidote is found in this artifact. **Titanite**
349 has 0.046 Al apfu. **Plagioclase** is albite with anorthite content 1-2 mole% (Figure 5b). **Chlorite** is
350 magnesian ($\text{Mg}\# = 0.66-0.67$) and relatively rich in Al (5.26-5.44 apfu). **Pumpellyite** composition
351 is pumpellyite-Al with $\text{Mg}\# = 0.94$.

352 4.3 Sample GR3-NIL.3

353 Sample GR3-NIL.3 (Figure 3C) has a dark to light green color. The texture is fine-grained and
354 massive. The mineral assemblage consists of jadeite (ca. 0.5 mm in size), omphacite (0.5 mm),
355 lawsonite (0.3 mm) phengite (0.1 mm), and titanite plus quartz in trace amounts and albitic
356 plagioclase and analcime (Figure 12; Table 2). Lawsonite and phengite show partial preferred
357 orientation (Figure 12a). Jadeite shows patchy and concentric zoning, locally oscillatory, and
358 replaces omphacite and lawsonite (Figure 12b-c). Interestingly, small grains of quartz occur as
359 inclusions within jadeite (Figure 12c). Plagioclase and analcime appear as discrete
360 inclusions/replacements within jadeite (Figure (Figure 12b). Textural analysis indicates that the
361 sequence of mineral growth is $\text{Omp}+\text{Lws}+\text{Ms}+\text{Ttn}(\text{Ab}?) \gg \text{Qtz}+\text{Jd} \gg \text{Ab}/\text{Anl}$. The presence of
362 lawsonite, the stability of $\text{Jd}+\text{Qtz}$ and the lack of epidote suggest relatively low temperature of
363 formation at high pressure.

364 **Jadeite** has varied composition (Figures 7 and 8). The total amounts of calculated Fe^{3+} and
365 Fe^{2+} are low, indicating jadeite-diopside solid solution with limited hedenbergite and aegirine
366 contents. This range is similar to the other described samples. Jadeite grains show concentric zoning
367 patterns (Figure 12b), with cores richer in Ca-Mg-Fe and rims richer in Na-Al, though the reverse is
368 also locally observed (see below). Furthermore, some crystals show oscillatory zoning with a
369 general increase in Ca-Mg-Fe towards the rims (Figure 12b). Finally, jadeite replacing omphacite
370 and lawsonite is zoned, with increasing Ca-Mg-Fe towards the outer zone of the replacement
371 (Figure 12d), but in general this type of jadeite has higher Ca-Mg-Fe (i.e., lower jd) contents than
372 matrix jadeite. These relationships indicate a complex history of jadeite growth which makes
373 resolving the general pattern of chemical evolution difficult. In this regard, however, it is important
374 to emphasize the trend towards more Ca-Mg-Fe rich compositions in the crystals with oscillatory
375 zoning, as is typical of Guatemalan jade (Harlow et al., 2011).

376 **Omphacite** occurs as inclusions within jadeite and as discrete crystals associated with
377 lawsonite (Figure 12b and d). Textural relations indicate omphacite replacement by jadeite (Figure
378 12b), suggesting that the inclusions represent omphacite grains replaced by jadeite. The
379 compositional range is relatively restricted (Figures 7 and 8) and indicates jadeite-diopside solid
380 solution with minor contribution of hedenbergite and aegirine components. It should be noted that
381 the composition of omphacite is not similar to that of previously described artifacts. The amounts of

382 jd and q components are higher and lower, respectively, which permit omphacite to plot fully within
383 the permitted region (at 500 °C) and results in the shrinking of the measured compositional gap
384 between jadeite and omphacite (Figure 8). This is consistent with a lower temperature of formation,
385 in agreement with the presence of lawsonite and lack of epidote in this artifact.

386 The composition range of **phengite** is relatively restricted, with Si up to 7.00 atoms apfu
387 (Figure 9). The amount of Tschermak (i.e., Si-Fe-Mg content) and paragonite (i.e., Na content)
388 substitutions are respectively higher and lower, than in phengite from the previously described
389 artifacts. This suggests lower temperature (and/or higher pressure) of formation. The amount of Ba
390 is low. **Plagioclase** is albite with 0.003 mole% anorthite content (Figure 5b). **Lawsonite** is almost
391 pure, with Fe = 0.01-0.02 apfu. **Titanite** has 0.04-0.05 Al apfu.

392 **4.4 Sample GR4-FIII.4**

393 Sample GR4-FIII.4 (Figure 3D) is green with locally darker patches and massive in texture.
394 The mineral assemblage consists of jadeite, titanite and hematite, with minor amounts of albite and
395 analcime (Table 2). We have not identified omphacite in this sample. The fine grained jadeite
396 matrix surrounds larger crystals of titanite and hematite (Figure 13). The crystals of jadeite are
397 locally replaced by late albite and analcime. The latter occupies clear interstitial positions along
398 cracks and within voids. Jadeite shows oscillatory zoning (Figure 13b), which zoning is truncated
399 by regions richer in Ca-Mg-Fe typically located adjacent to voids/fractures filled with analcime
400 (Figure 13b). The deduced sequence of crystal growth is Jd+Hem+Ttn >> Ca-Fe-Mg rich
401 Jd/Ab/Anl.

402 The range in composition of **jadeite** is relatively restricted (Figures 7 and 8). This
403 composition differs from that of other artifacts in having higher aeg content, consistent with the
404 presence of the Fe³⁺ saturated phase hematite and, henceforth, high fO₂ conditions. Also, jadeite
405 from this artifact reaches almost pure jadeite composition.

406 **Hematite** is nearly pure (Ti = 0.00, Mn < 0.003 apfu). **Plagioclase** is albite (1 mole %
407 anorthite; Figure 5d). **Titanite** has 0.08 Al apfu. Analcime is almost pure (Ca = 0.01 apfu).

408 **5 Discussion and Conclusions**

409 **5.1 Source region of the artifacts**

410 The detailed discussion on the composition of the jadeite artifacts recovered from Golden
411 Rock attests to the need to conduct multi-faceted characterization studies of this raw material in
412 order to come to terms with the geological signatures that will allow us to pinpoint its most

413 probable provenance in the Caribbean and elsewhere. To our knowledge, the work of Harlow et al.
414 (2006) on the late Saladoid artifacts from the island of Antigua is the only published in-depth
415 mineralogical-petrological study of jadeitite artifacts within the Caribbean that has provided
416 comparable results to the ones hereby presented. The other studies on jadeitite identification and
417 characterization in the Antilles (Hardy, 2008; Johnson, 1980; Rodríguez Ramos, 2007) were
418 constrained to non-destructive X-ray diffraction techniques, which are basically restricted to
419 mineral identification and limit the range of minerals to be identified to only those present in
420 sufficient concentrations in the jadeitite (see for example Harlow et al., 2006, table 2). Given the
421 subtle differences in mineralogical composition between the different jadeitite sources discussed in
422 this paper, it is obvious that these X-ray diffraction results are not as useful in pinpointing the exact
423 provenance of jadeitites.

424 From a mineralogical-petrological point of view, the Golden Rock artifacts on nearby St.
425 Eustatius are similar to their Antiguan equivalents. Samples containing phengite+biotite/phlogopite,
426 lawsonite, or quartz are common at both sites (Table 3). However, there are some differences as
427 well, which include the presence of paragonite or glaucophane in some of the Antiguan artifacts,
428 which have not been identified within the St. Eustatius assemblage, as well as abundant hematite in
429 one St Eustatius specimen, absent from the Antigua artifacts. However, these differences may be
430 the result of sample bias, taking into account the limited number of artifacts analyzed and the
431 considerable variation that can exist between different jadeitite pieces from a single locality.
432 Despite these differences, it is likely that jadeitite artifacts from both archaeological sites were
433 extracted from the same source region.

434 On evaluating the provenance of the Antiguan artifacts, Harlow et al. (2006) indicated that
435 they matched with jadeitite from the South Motagua Fault serpentinite mélanges of Carrizal Grande,
436 La Ceiba and Ensenada, particularly given the presence of quartz, phengite+biotite/phlogopite,
437 lawsonite, glaucophane and the lack of preiswerkite (Harlow et al., 2006, 2011; Table 3). Harlow et
438 al. (2006), however, also reported important mismatches with this source, such as the presence of
439 paragonite and (c)zoisite/epidote+quartz±lawsonite in the Antiguan artifacts that have not been
440 described in jadeite from the South Motagua mélange. Paragonite and (c)zoisite/epidote are,
441 however, abundant in jadeitite from the North Motagua mélange, but quartz and lawsonite are
442 absent. Due to these dissimilarities, Harlow et al. (2006) concluded that other unknown sources in
443 the Caribbean realm could not be ruled out, and specifically suggested the serpentinite mélanges
444 from Cuba and Dominican Republic as possible alternatives.

445 With the discoveries of jadeitite occurrences on those two islands, a comparison with the
446 results from the study of the Antiguan material can now be made. As evident in Table 3, key

447 mineralogical characteristics do not match with the material from Cuba or the Dominican Republic,
448 so these were not the source of the Antiguan artifacts. Even if paragonite and
449 (c)zoisite/epidote/allanite±quartz, also common in the Cuban material, are characteristics of the
450 Antiguan artifacts, the presence of lawsonite and glaucophane in the artifacts is a significant
451 mismatch. These two minerals have not been observed in any studied samples from the Cuban
452 source. The jadeitites from the Rio San Juan deposit, on the contrary, commonly bear abundant
453 glaucophane, but paragonite and biotite/phlogopite have not been described. Furthermore, lawsonite
454 is common in this deposit, but not in jadeite s.s. but in jadeite-lawsonite quartzite and lawsonite-
455 quartzite, two rather different rock types than that of the Antiguan artifacts.

456 The source assignment of the St. Eustatius artifacts suffers from similar uncertainties. Since
457 the studied artifacts represent exchanged items that may come from different sources, we have
458 defined three groups based on mineralogical assemblages: Group 1 (GR24-B-III.1 and GR24-J-
459 IV.05) consists of samples bearing phlogopite (plus phengite and epidote but minus lawsonite and
460 quartz); Group 2 (GR3-NII.3) represents the sample bearing lawsonite (plus phengite and quartz but
461 minus phlogopite and epidote); and Group 3 (GR4-FIII.4) is defined by containing significant
462 hematite. Significantly, we have not identified paragonite or glaucophane in any of the artifacts. In
463 general terms the different groups can be associated with different known sources. Group 1 exhibits
464 strongest similarity with the North Motagua and Sierra del Convento deposits, while group 2 fits the
465 South Motagua and Río San Juan deposits best, and group 3 matches up the North Motagua deposit,
466 particularly that of Cerro Colorado (Harlow et al., 2011). Major element composition of the
467 minerals does not add much in discriminating the sources, as illustrated by the composition of
468 clinopyroxenes and phengite. The published analyses of these minerals have been compared with
469 those of the studied artifacts (Figure 14). The micas compare well with the South Motagua source,
470 but the scarcity of published analyses for the Sierra del Convento and Rio San Juan mélanges
471 prevent definitive conclusions. The case for pyroxenes is similar, though the uncertainties in
472 calculated Fe^{3+} and aegirine contents prevent further considerations. Hence, at present,
473 mineralogical assemblages constitute the best proxy for evaluating provenance (Table 3).

474 In this regard, even if the Sierra del Convento deposit could be raised as a potential source of
475 the phlogopite-bearing artifacts, the lack of lawsonite in this deposit (but present in Group 2) is a
476 major mismatch. Similarly, the Dominican deposit seems an unlikely source due to the lack of
477 phlogopite/biotite (present in Group 1), the common presence of glaucophane (absent in the
478 artifacts), and the presence of lawsonite only in quartz-rich rocks. Also, hematite-bearing samples
479 are lacking in these two deposits. There is also discrepancy with the North Motagua mélange, such
480 as the lack of paragonite in group 1, but this mineral is only present in a portion of the North

481 Motagua samples and the discrepancy may therefore be insignificant. Only the South and North
482 Motagua mélanges clearly stand out as likely sources for groups 2 and 3, respectively.

483 Consequently, it can be concluded that the Motagua Valley (North and South Motagua
484 mélanges) is the most likely source for artifact groups 2 and 3, and probably also group 1. It is
485 plausible that both Guatemalan deposits supplied the trade goods that arrived to the Golden Rock
486 site. This conclusion agrees with the South Motagua source assignment of the Antigua jadeitite
487 artifacts by Harlow et al. (2006), though our conclusion of multiple Guatemalan sources for Golden
488 Rock seems to make more sense because no major political/geographical barrier that may have
489 hampered trade divided the North and South Motagua Valley deposits in pre-Columbian times.
490 However, additional deposits yet to be discovered are possible (Jamaica, Nicaragua, Colombia and
491 Venezuela; Figure 1), at least for the Antigua artifacts which do not fully match with the known
492 sources. Likewise, new finds and increased diversity within the known sources is another
493 possibility, as the source areas and diversity in Guatemala alone have grown since the work on the
494 Antigua artifacts in 2005; the picture is likely to change.

495 **5.2 Archaeological implications**

496 The jadeitite axes from Golden Rock clearly suggest the existence of interaction webs that led
497 to the import of this important raw material into the Lesser Antilles in contexts post-dating AD 600.
498 This evidence corresponds to that recovered from contemporaneous sites in Puerto Rico (e.g., Río
499 Tanamá and Tecla II) and other islands in the Lesser Antilles (e.g., Spring Bay 1a on Saba, Elliot's
500 on Antigua) where this type of material has also been documented (Rodríguez Ramos, 2007, 2011).
501 The higher frequencies of these jadeitite artifacts in the northern Antilles when compared to the
502 southern islands, indicates the unlikelihood that these circulated from northeastern South America
503 via the Lesser Antilles. This opens the door for one or more of the following scenarios: a) the direct
504 trans-Caribbean circulation of axes through interaction spheres articulated between these
505 northernmost islands and the Isthmo-Colombian area, to where jadeitites were moved south from
506 their original source in Guatemala (Harlow, 1993); b) that these materials came from one of the
507 recently identified sources in the Dominican Republic or Cuba, or; c) that materials come from a yet
508 unidentified source in either the Greater Antilles and Central or South America.

509 The petrological evidence hereby presented tends to support the first possibility, as it
510 indicates that the Motagua River Valley is the most likely source for these materials, as was also
511 suggested in previous research on materials from Antigua (Harlow et al., 2006), Puerto Rico
512 (Harlow, 2006, unpublished, cited in Rodríguez Ramos, 2007) and the Virgin Islands (Harlow,
513 2008, unpublished, cited in Hardy, 2008). The presence of Motaguan jadeitite in these islands might
514 serve as evidence of the participation of its inhabitants in these long-distance trans-Caribbean

515 networks that seem to have been configured since the early peopling of the islands. Although the
516 evidence of jadeitite pieces in the Antilles pre-dating AD 600 is scant, being limited to several
517 petaloid celts possibly related to pre-AD 600 occupation phases of the multi-component Tecla 1 and
518 Punta Candellero sites (Harlow 2006, unpublished, cited in Rodríguez Ramos 2007), there is clear
519 evidence of other continental materials coming into the northern Antilles from the continental
520 Caribbean since early on in the occupation of the Antilles that reaffirm the possibility for a
521 continental origin for these jadeitite artifacts. The presence of these continental products in the
522 northern Antilles initially involved the circulation of cultivars such as maize, manioc, and sweet
523 potatoes during the Archaic Age and later. During the Early Ceramic Age (400 BC – AD 800), it
524 entailed the movement of jaguar, peccary, and tapir tooth pendants, guanín (i.e., tumbaga or gold
525 copper alloys), nephrite adornments, and possible other products that have clearly an extra-
526 Antillean derivation (Rodríguez Ramos, 2010, 2011; Rodríguez Ramos and Pagán Jiménez, 2006).
527 The iconographic resemblances between personal adornments documented in the northern Antilles
528 and those observed in the Isthmo-Colombian region, which conforms the southernmost distribution
529 of Motaguan jadeitite documented thus far in the American continents, has led to the suggestion of
530 the existence of an Isthmo-Antillean interaction sphere that started to take shape since the earliest
531 occupation of the islands (Rodríguez Ramos, 2013). The jadeitite distribution networks were
532 inserted into other Antillean webs through which stone celts were being moved, particularly grey-
533 green tuffaceous mudstone from St. Martin and silicified tuff from eastern Puerto Rico
534 (Knippenberg, 2004, 2006; Rodríguez Ramos, 2001). These webs also included the circulation of
535 porphyry *cemí* three pointers from St. Martin (Knippenberg, 2004, 2006). These webs must initially
536 have been a very important means by which the earliest Ceramic Age settlers secured their survival
537 in a new and sparsely populated islands setting (Hofman et al., 2007, 2011; Knippenberg, 2006).

538 Although the evidence presented here pinpoints the Motagua area as the exclusive source for
539 the jadeitites documented in St. Eustatius, Antigua, Puerto Rico and the Virgin Islands, this line of
540 research will benefit immensely with more complete characterization of the Dominican and Cuban
541 sources, which is underway (e.g., Cárdenas-Párraga et al., 2012; García-Casco et al., 2009; Schertl
542 et al., 2012; Hertwig et al., 2012), and the search for new sources in the Circum-Caribbean region.
543 However, regardless of its source, the jadeitite material at Golden Rock represented a true exotic for
544 its inhabitants, and it likely had one of the most distant origins of all lithics found at the site, much
545 further than the grey-green mudstone celts and calcirudite conglomerate *cemí* three-pointer stones
546 from neighboring St. Martin or flint material and tools from Antigua (Knippenberg, 2006). Despite
547 its distant provenance, jadeitite axes were primarily used and valued as tools at Golden Rock, as
548 well as at many other sites in which these have been documented. The relative hardness and
549 toughness of the rock compared to the other available rock types in the region must have been the

550 main reason for this. Still examples exist from undated and unpublished contexts that show that in
551 some instances it also fulfilled ceremonial or ritual functions, which may primarily be associated
552 with its' brilliant lustre, green colour, and exotic nature (see also Berman, 2011; Rodríguez Ramos,
553 2007, 2011).

554 Its presence at Golden Rock furthermore shows that, despite the disappearance of the
555 extensive Early Ceramic Age bead and pendant network somewhere between AD 400 and 600
556 (Knippenberg, 2006), other materials such as jadeitite became important in long-distance
557 interactions. It possibly also signifies that networks were re-directed. This likely was the result of a
558 combination of factors, including population growth, cultural diversification, and migration of
559 groups to the Lesser as well as the Greater Antilles. These processes shifted the social landscape
560 and opened up new territories, including previously unused raw material sources.

561 Following this early presence at Golden Rock during the late phase of the Early Ceramic Age,
562 jadeitite use has been firmly attested at different well dated Late Ceramic Age (AD 800 – 1500)
563 sites, such as for example Anse à la Gourde on Guadeloupe, and El Cabo and Playa Grande in the
564 Dominican Republic, Tecla II and Rio Tanamá in Puerto Rico and other sites in the Virgin Islands
565 and the Grenadines. This evidence suggests that the use of jadeitite had become a structural element
566 in Caribbean Amerindian culture until the late pre-Columbian occupation of the islands.

567 This exchange involving jadeitite, along with other materials and information, was part of
568 complex multi-vectorial networks that shifted in configuration and directionality throughout the
569 pre-colonial history of the Antilles (see also Hofman and Hoogland, 2011; Hofman et al., 2007,
570 2011). Hopefully, this study provides a step forward towards a better understanding of the millenary
571 engagements that took place within and between the insular and the continental Caribbean.

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699 95, pages 185–187.

700 **8 Figure Captions**

701 Figure 1. Location of known jadeitite deposits in the Caribbean region, of studied archaeological jadeitite from Antigua
702 (Harlow et al., 2006) and St. Eustatius (this work), of ultramafic bodies (high-pressure or not), and of potential regions
703 where jade deposits (related to high-pressure ultramafic rocks) are possible (see the electronic version of the article for a
704 color version of this figure).

705 Figure 2. Golden Rock Site lay-out showing the location of the jadeitite artefacts discussed in this paper in relation to
706 the refuse midden area and house plans belonging to the different occupation phases at the site (modified after Versteeg
707 and Schinkel, 1992, fig. 161, p.181). Numbers of the artefacts refer to numbers in Table 1. The inset shows the location
708 of the site in the island (see the electronic version of the article for a color version of this figure).

709 Figure 3. Photographs of the studied artifacts from Golden Rock site (St. Eustatius). The scale shown in the photographs
710 is given in cm. A) Fragment of axe, proximal-lateral part (sample GR24-B-III.1). B) Edge flake (sample GR24-J-
711 IV.05). C) Fragment of axe (sample GR3-NII.3). D) Edge flake (sample GR4-FIII.4). See the electronic version of the
712 article for a color version of this figure.

713 Figure 4. Textures and mineral associations of artifact GR24-B-III.1. A) and B) Optical images (crossed polars) of the
714 sample showing large crystals of jadeite and phengite. C) and D) BSE images. Note omphacite included in jadeite and
715 the clots of omphacite, locally associated with phengite. E) and F) Close up BSE images (see location in C) showing
716 corroded omphacite within jadeite and the patchy zoning of jadeite (darker color denotes Na-Al richer composition).
717 See the electronic version of the article for a color version of this figure.

718 Figure 5. ACFN "deluxe" diagrams for the analyzed minerals of the studied artifacts calculated with *CSpace* (Torres
719 Roldán et al., 2000). The minerals are projected from the phases and exchange vectors indicated in A. Representative
720 end-members of the solid solutions are indicated (lower case font). See the electronic version of the article for a color
721 version of this figure.

722 Figure 6. XR maps of artifact GR24-B-III.1. The color scale bar represents relative element concentration (counts; red
723 and purple denote higher and lower concentration, respectively). The horizontal scale bar is 1 mm. See the electronic
724 version of the article for a color version of this figure.

725 Figure 7. Binary diagrams showing the composition of analyzed clinopyroxenes expressed in terms of cations per
726 formula unit (6 oxygens), with indication of significant end-members and exchange vectors. See the electronic version
727 of the article for a color version of this figure.

728 Fig. 8. Composition of jadeite and omphacite plotted in the ternary jadeite-q-aegirine diagram with the clinopyroxene
729 classification scheme of Morimoto et al. (1988; names in italics). Also shown are the phase relations of clinopyroxene
730 (denoted by space groups *C2/c* and *P/2n*) calculated by Green et al. (2007) at 500 °C in the system jadeite-diopside-
731 aegirine. Note that jadeite and omphacite with higher q component plot within the omphacite-jadeite and omphacite-
732 diopside solvus (forbidden regions) as a likely consequence of the lack of consideration of hedenbergite in the
733 calculated phase relations and/or different temperature of formation. Solid lines indicate varying aegirine/(q+aegirine)
734 ratios. See the electronic version of the article for a color version of this figure.

735 Figure 9. Binary diagrams showing the composition of K-white micas. Representative end-members and substitution
736 vectors are shown (note that, for convenience, the "celsian" vector has been drawn to intersect the phengite end-
737 member). See the electronic version of the article for a color version of this figure.

738 Figure 10. Phlogopite composition plotted in the "ideal plane of biotite" (Deer et al., 1962; Guidotti, 1984)
739 Representative end-members are shown. See the electronic version of the article for a color version of this figure.

740 Figure 11. Textures and mineral associations of artifact GR24-J-IV.05. A) Optical (crossed polars) and B) BSE images
741 showing the nodular structure and Omp+Ep+Jd intergrowths. C) Optical (crossed polars) and D) BSE images showing
742 details of the intergrowth texture. For convenience, the BSE image is presented in a color scale (red color indicates
743 higher average atomic weight; red: epidote; green: omphacite; blue: jadeite). See the electronic version of the article for
744 a color version of this figure.

745 Figure 12. Textures and mineral associations of artifact GR3-NII.3. A) Optical (crossed polars) image showing the
746 general texture of the sample. B) BSE image showing the association of lawsonite+phengite and the chemical zoning of
747 jadeite (i.e., cores richer in Ca-Mg-Fe, rims richer in Na-Al and local oscillatory zoning). Note that lawsonite is
748 replaced by jadeite. C) BSE image showing quartz inclusions within jadeite. Lawsonite is replaced by jadeite. D)
749 Replacement of omphacite and lawsonite by jadeite. Note that jadeite replacements are richer in Ca-Mg-Fe than
750 jadeite in the matrix. See the electronic version of the article for a color version of this figure.

751 Figure 13. Textures and mineral associations of artifact GR4-FIII.4. A) and B) Optical (crossed polars) and BSE images
752 showing the texture of the sample. Note oscillatory zoning of jadeite crystals and the Ca-Mg-Fe richer composition of
753 jadeite adjacent to analcime crystals grown in voids/fractures. See the electronic version of the article for a color version
754 of this figure.

755 Figure 14. Comparison of phengite (left) and clinopyroxene (right) composition from St. Eustatius jade artifacts with
756 Antiguan jade artifacts (Harlow et al., 2006) and jadeitites from the mélanges of North and South Motagua (Harlow et
757 al., 2006), the Sierra del Convento (Cárdenas-Párraga et al., 2012) and the Río San Juan (Schertl et al., 2012). For
758 clarity, only the clinopyroxenes from the South Motagua mélange have been plotted. See the electronic version of the
759 article for a color version of this figure.

1 **Pre-Columbian jadeitite artifacts from the Golden Rock Site, St.**
 2 **Eustatius, Lesser Antilles, with special reference to jadeitite**
 3 **artifacts from Elliot's, Antigua: Implications for potential source**
 4 **regions and long-distance exchange networks in the Greater**
 5 **Caribbean**

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18 **Tables**

19 Table 1. Context and dimensions of Golden Rock jadeitite artefacts. Level is below surface; MD=maximum dimension;
 20 L=length, W=width; Th=thickness (all in mm); W=weight (in g). All artifacts are considered to date AD 600 and 825-
 21 890 (see text).

NR.	Number	1x1 m	level	Artifact type	MD	L	W	Th	W	Context
1	GR 24/B/III.01	24 B	20-30 cm	axe fragment, proximal-lateral	43	43	26	25	26,4	Refuse deposit; test-unit
2	GR 24/J/IV.05	24 J	30-40 cm	edge flake	25	25	25	6	4,2	Refuse deposit; test-unit
3	GR 3/N/II.03	3 N	10-20 cm	axe fragment, medial-lateral	29	29	15	9	3,5	Refuse deposit; test-unit
4	GR 4/F/III.04	4 F	20-30 cm	edge flake	36	29	34	6	6,8	Refuse deposit; test-unit
	GR 15/K/I.06	15 K	0-10 cm	edge flake	26	25	26	4	3,3	Refuse deposit; test-unit
	GR 15/G/I.01	15 G	0-10 cm	flake from axe	26	26	18	9	3,4	Refuse deposit; test-unit
	GR 20/B/III.01	20 B	20-30 cm	axe fragment, medial-lateral	39	39	31	26	34,1	Refuse deposit; test-unit

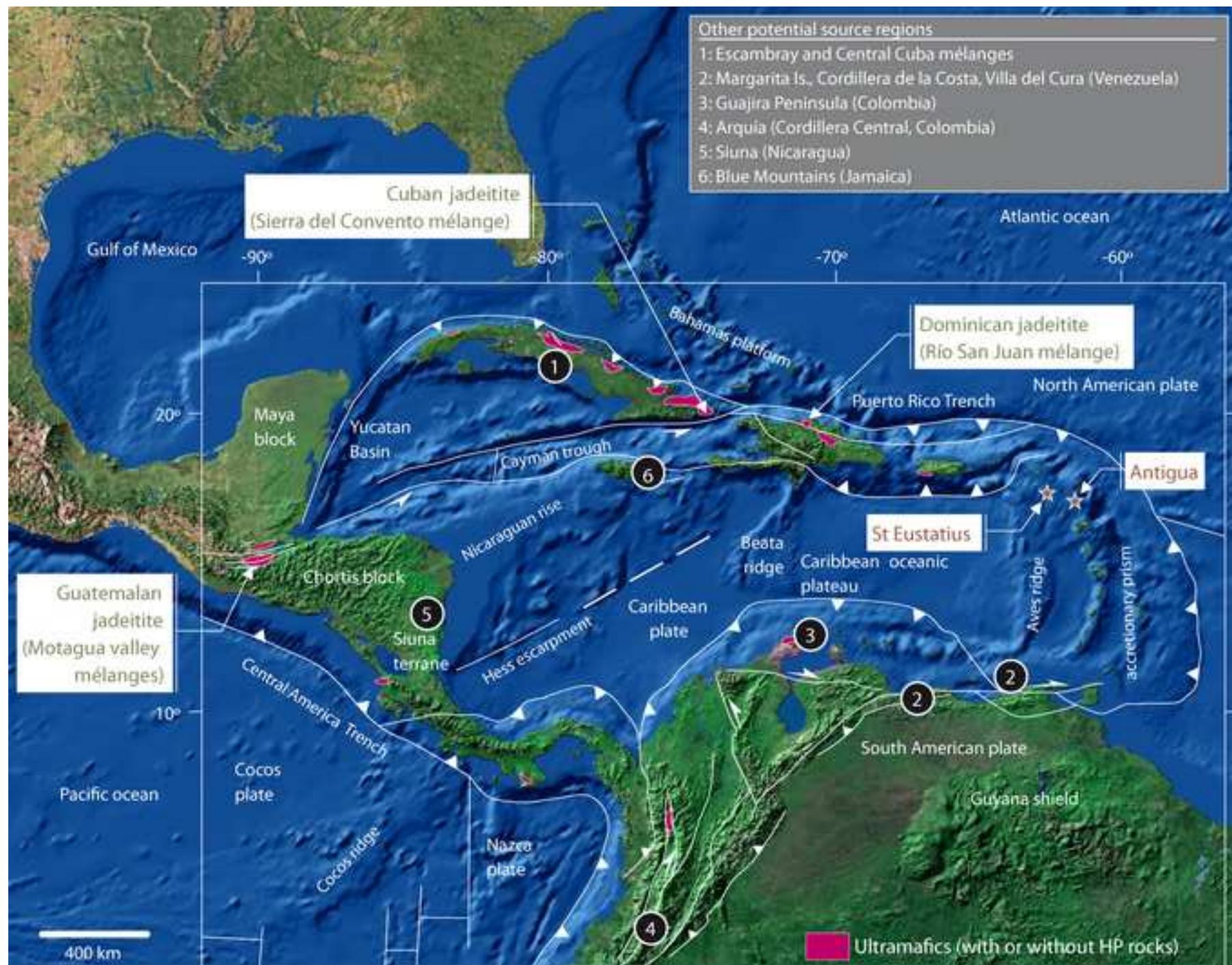
22 Table 2. Mineral assemblages of studied samples. Symbol "x" denotes the relative abundance of minerals. Symbol "-"
 23 indicates minor/trace amounts.

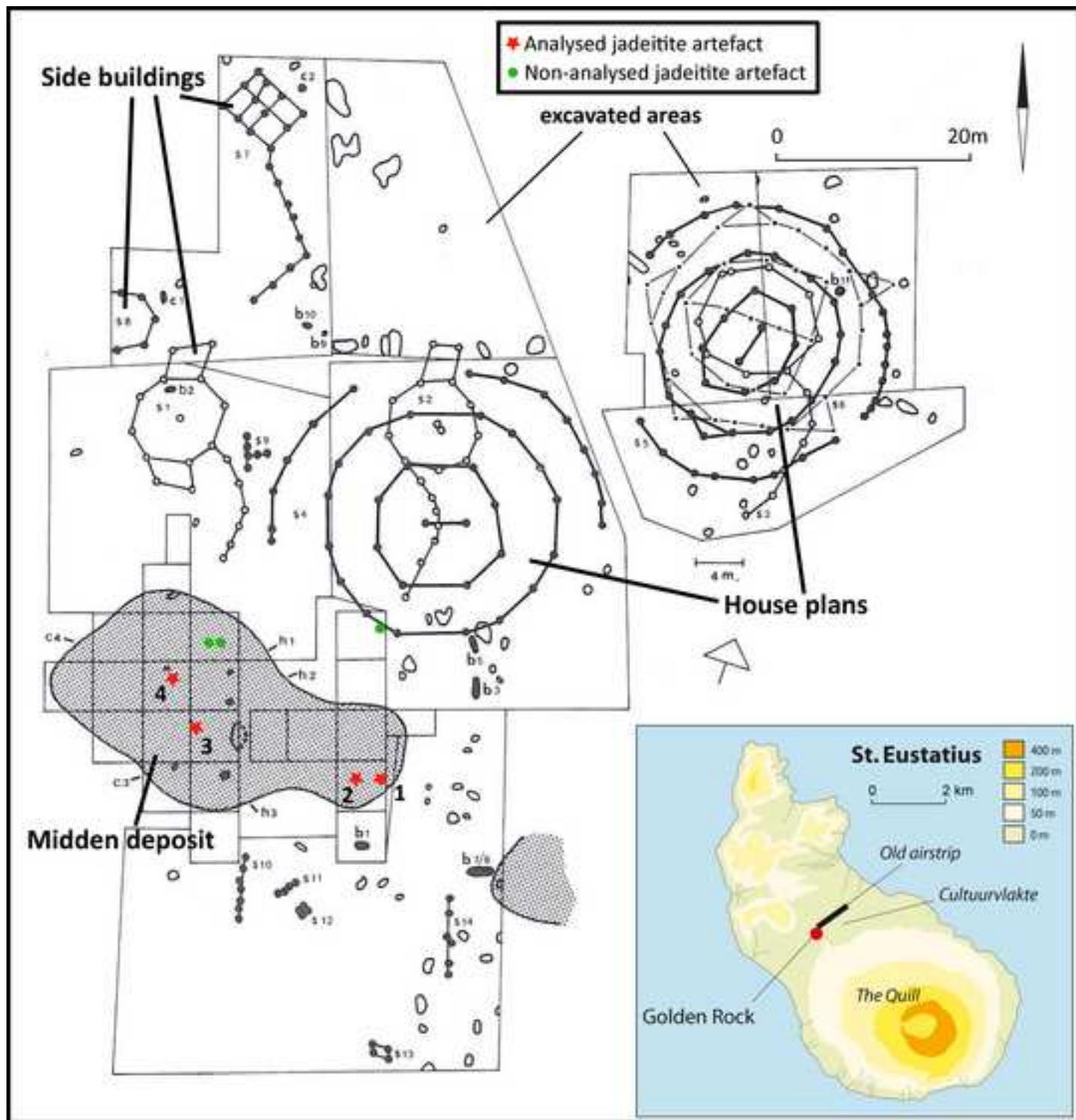
	GR24-B-III.1	GR24-J-IV.05	GR3-NII.3	GR4-FIII.4
Jadeite (Jd)	xxx	xxx	xxx	Xxx
Omphacite (Omp)	xx	xx	xx	
Albite (Ab)	-	-	-	-
Analcime (Anl)	-			-
Phengite (Ph)	xx (Ba-Ph)	x	x	
Phlogopite (Phl)	x	x		
Epidote(Ep)/(clino)Zoisite (Czo)	x Ep	xx (Czo)		
Lawsonite (Lws)			x	
Titanite (Ttn)	x	x	-	x
Quartz (Qz)			-	
Chlorite (Chl)		-		
Pumpellyite (Pmp)		-		
Hematite (Hem)				x

24 Table 3. Mineralogical comparison of studied artifacts with jadeite artifacts from Antigua (Harlow et al., 2006) and
 25 with jadeite from known Caribbean sources in the North and South Motagua mélanges, Guatemala (Harlow et al.,
 26 2006; 2011), Sierra del Convento mélange, eastern Cuba (Cárdenas-Párraga et al., 2010, 2012; García-Casco et al.,
 27 2009) and Río San Juan mélange, Dominican Republic (Schertl et al., 2012). The shaded cells denote critical mineral
 28 for provenance analysis. Note that all minerals indicated for the different regions do not necessarily coexist in
 29 individual samples.

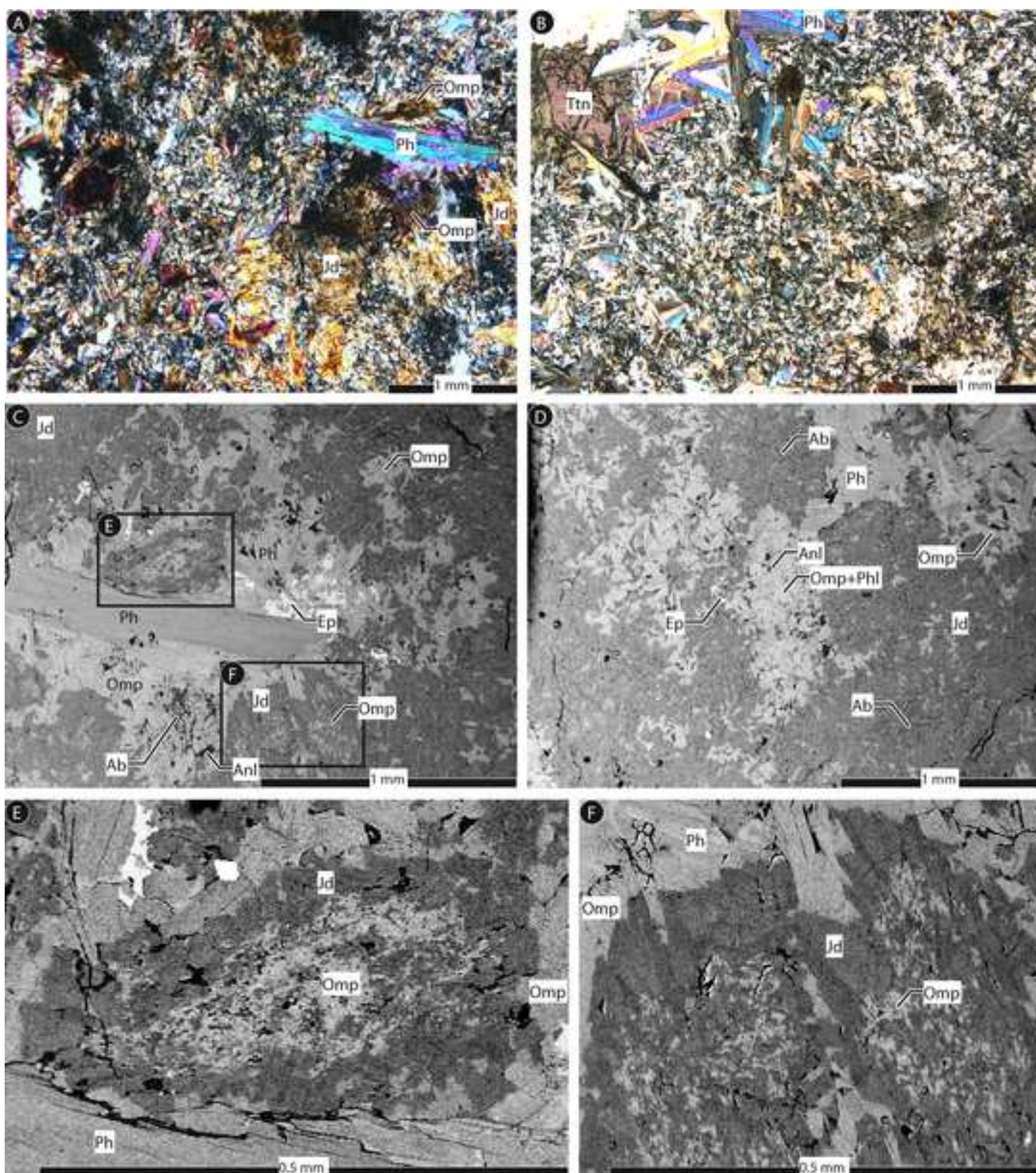
Sample	Jd	Omp	Ph	Phl	Qz	Pa	Ttn	Lws	Ep	Gl	Pl	Ap	Chl	Pmp	Anl
GR24-B-III.1	x	x	x	x			x		x		x	X			x
GR24-J-IV.05	x	x	x	x			x		x		x		x	X	
GR3-NII.3	x	x	x		x		x	x			x				
GR4-FIII.4	x						x				x				x
From South Motagua source? (Harlow et al., 2006)															
Antigua	x	x	x	x	x	x (local)	x	x (local)	x	x (local)	x	X		X	x
Known Sources															
NM (Gua)	x	x	x	x		x	x		x		x	X	x	X	x
SM (Gua)	x	x	x	x	x		x	x	x	x	x	X	x	X	x
RSJ (DR)	x	x	x		x		x	x (local)	x	x (abun.)	x	X	x		x (local)
SC (Cuba)	x	x	x	x	x	x	x		x		x	X	x	X	x

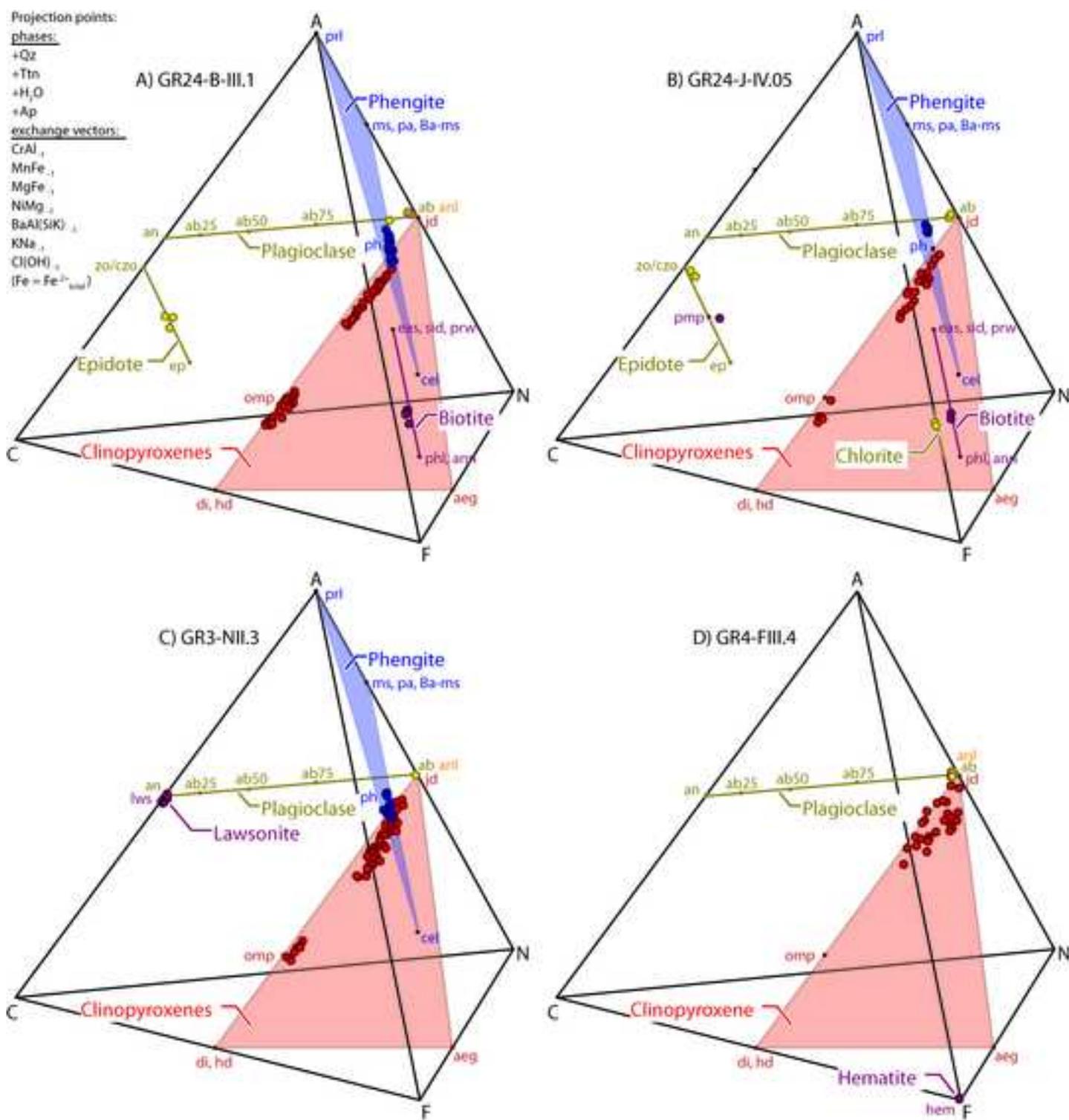
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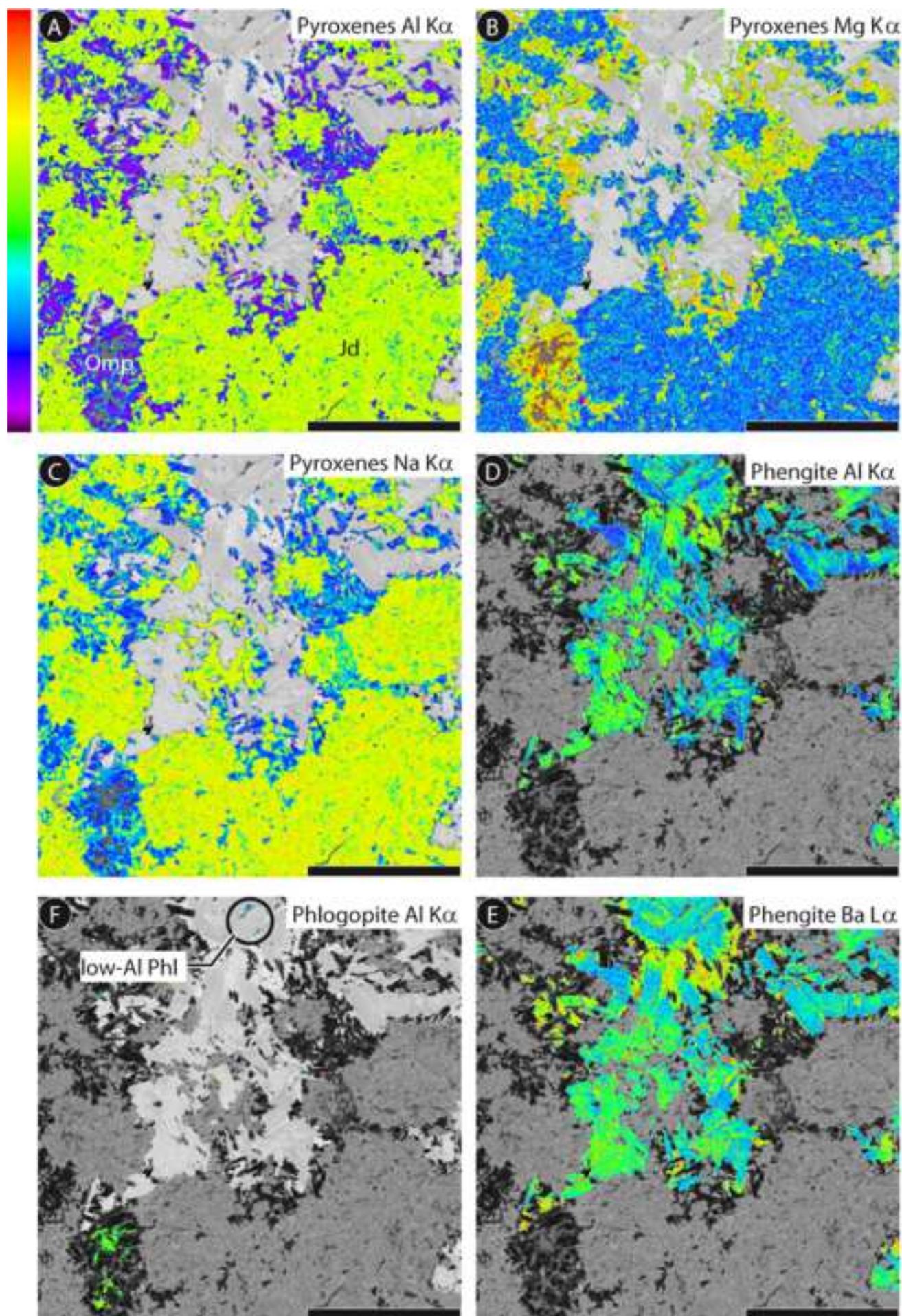




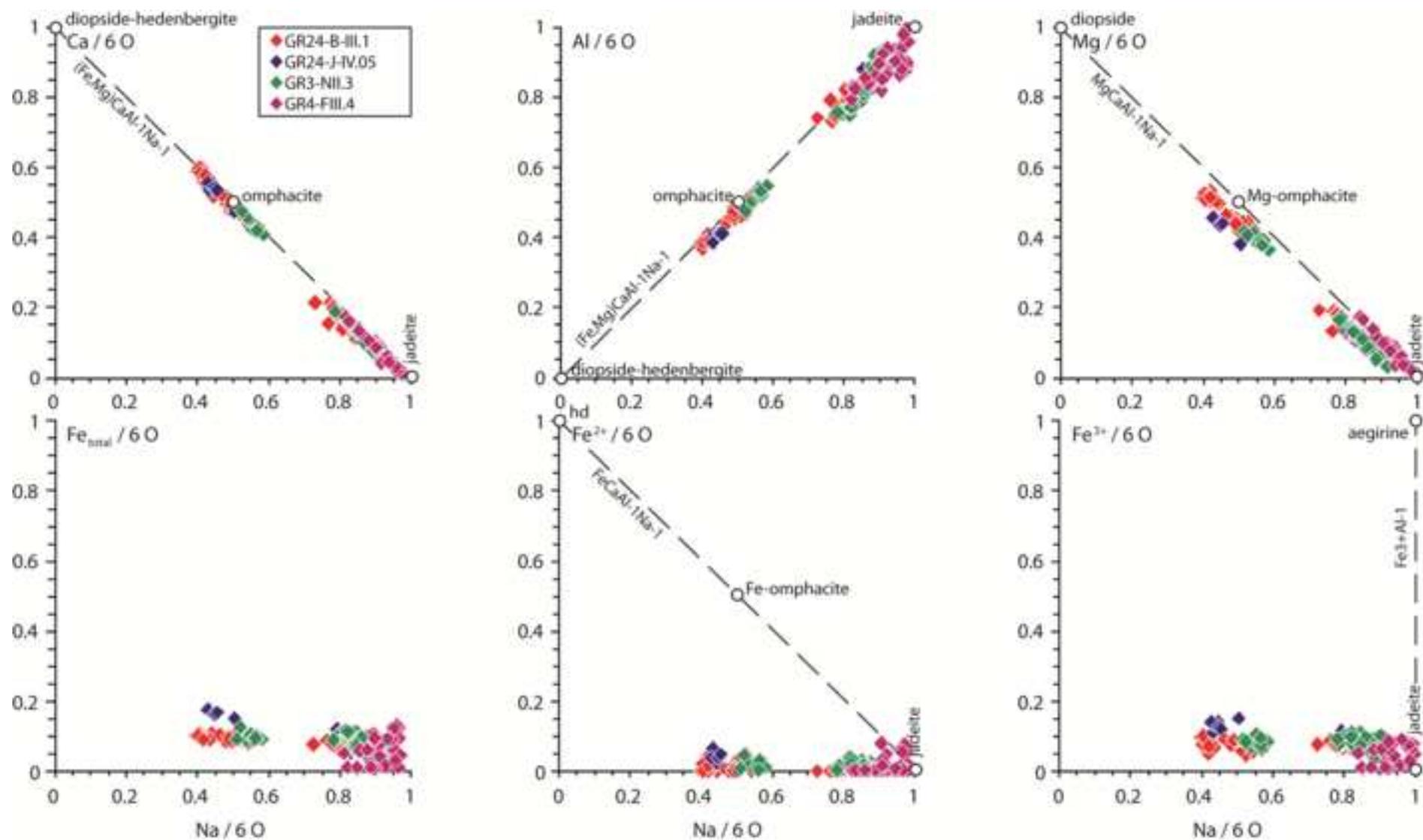




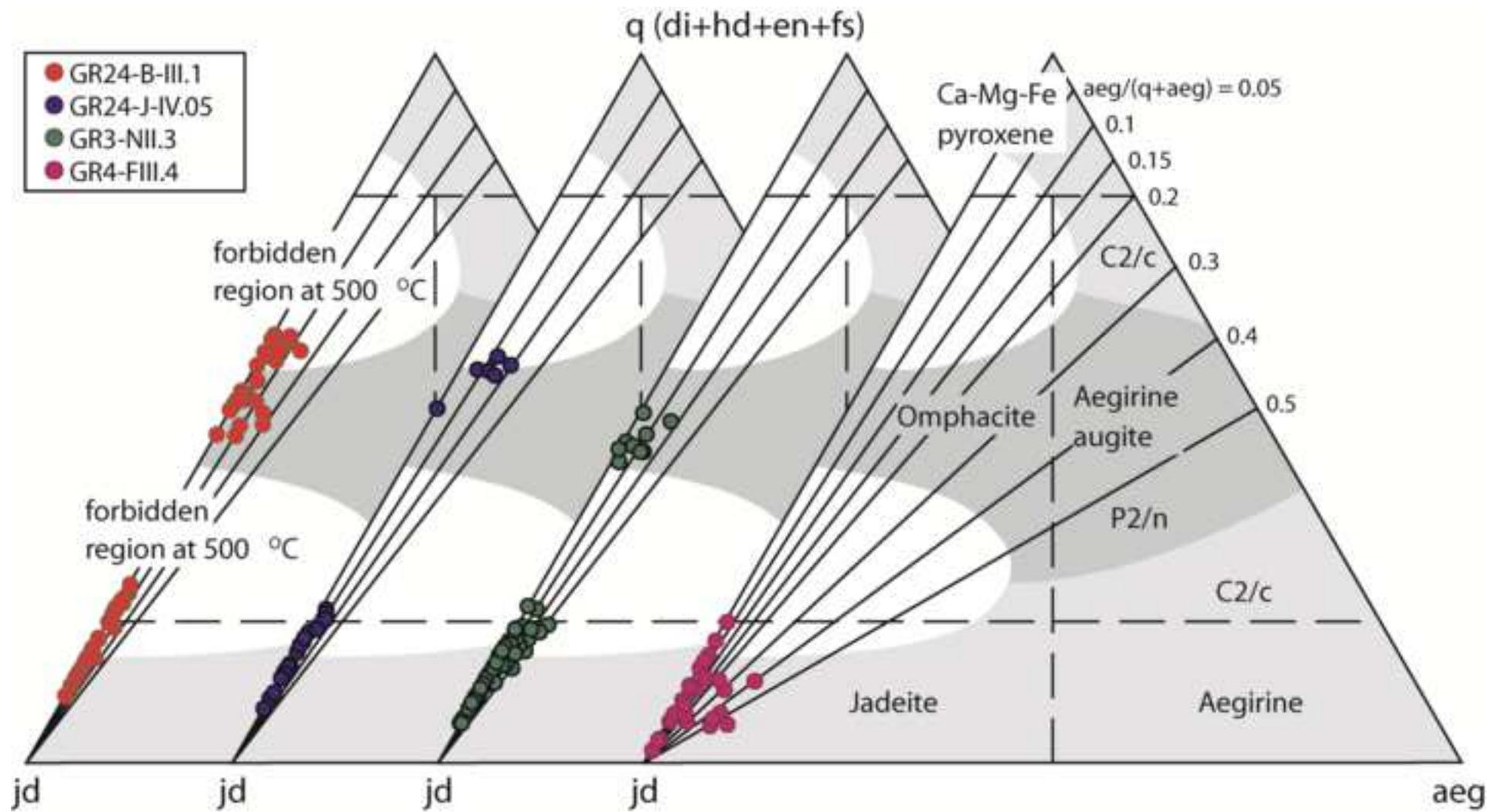


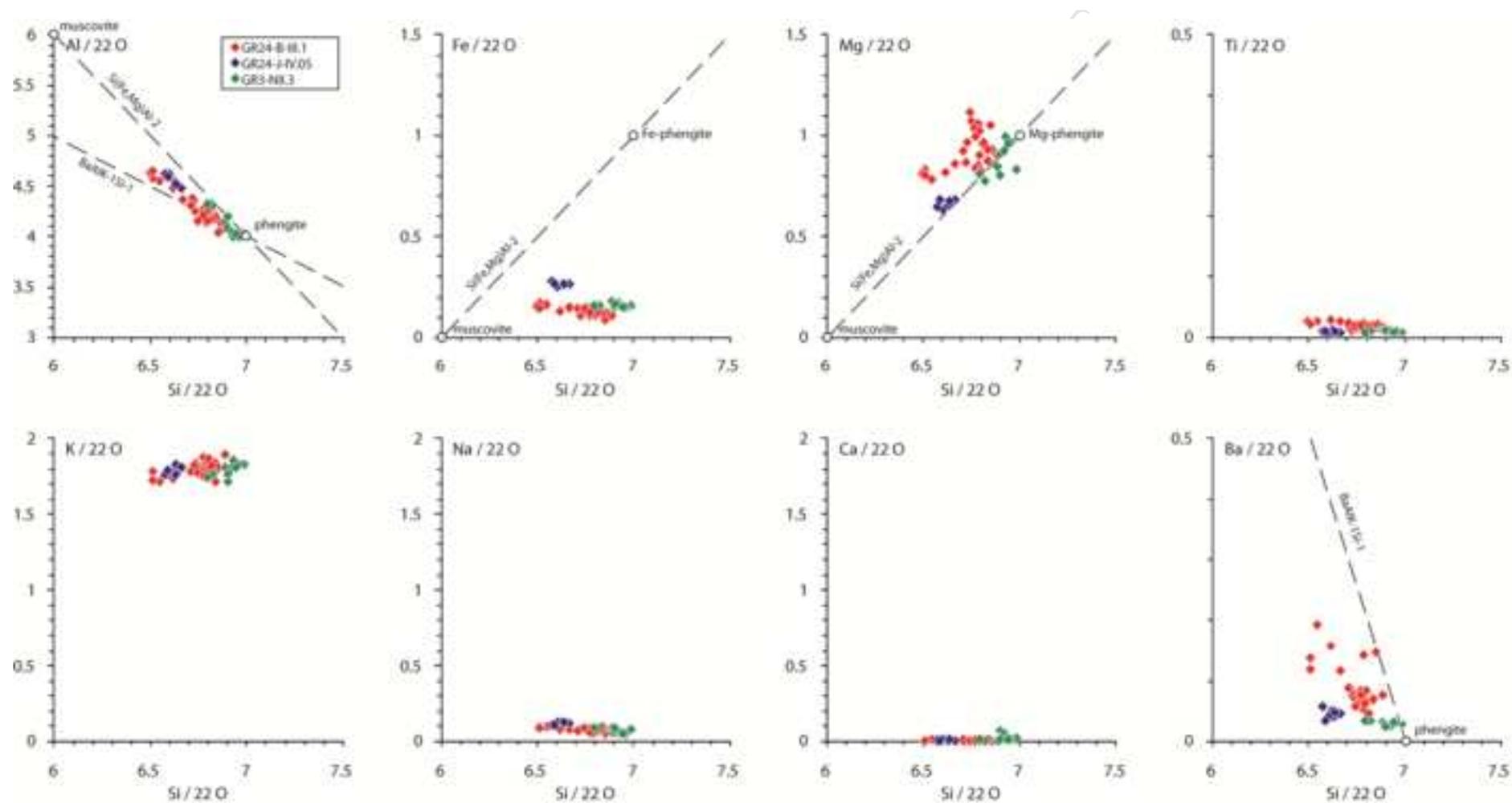


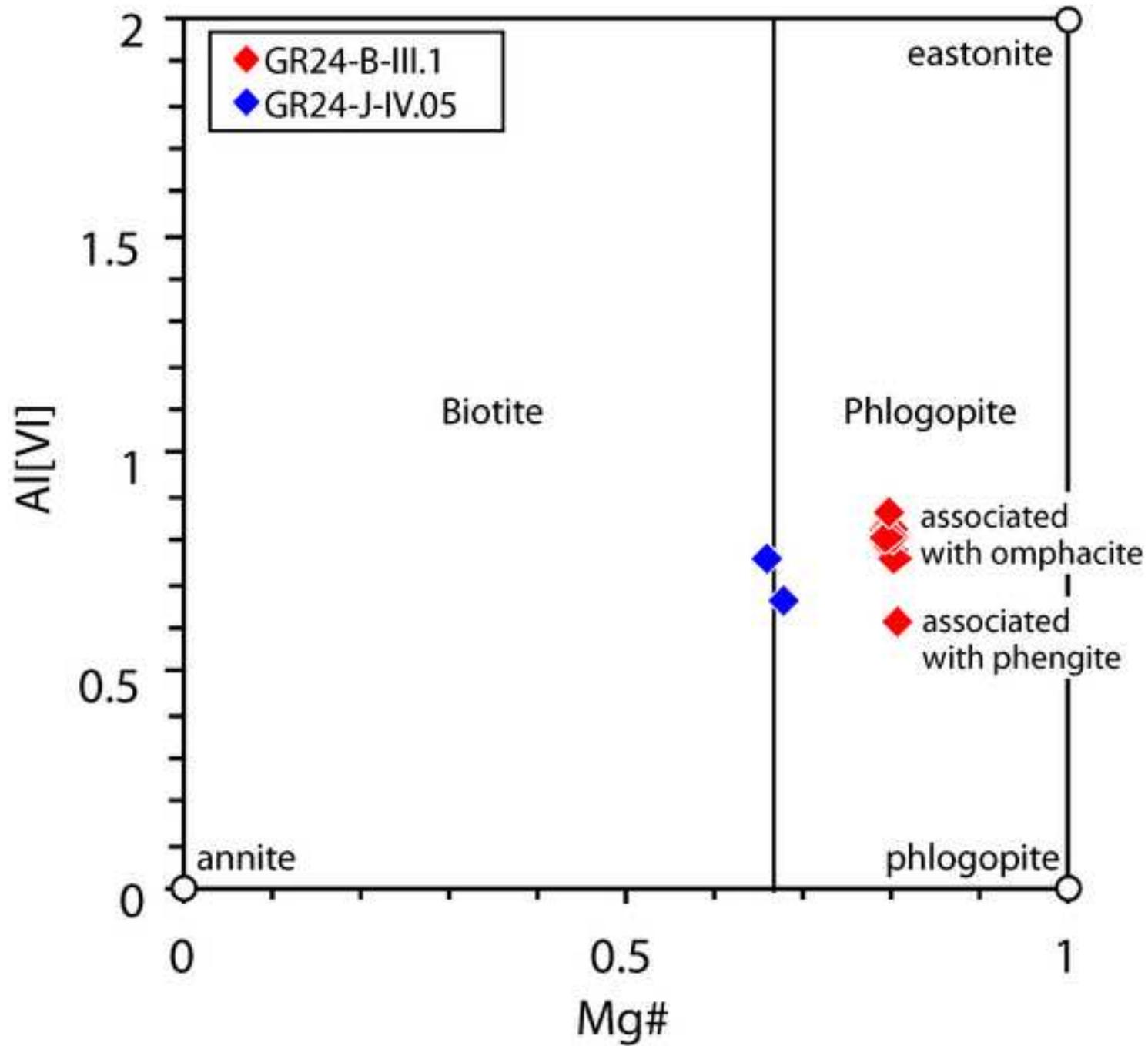
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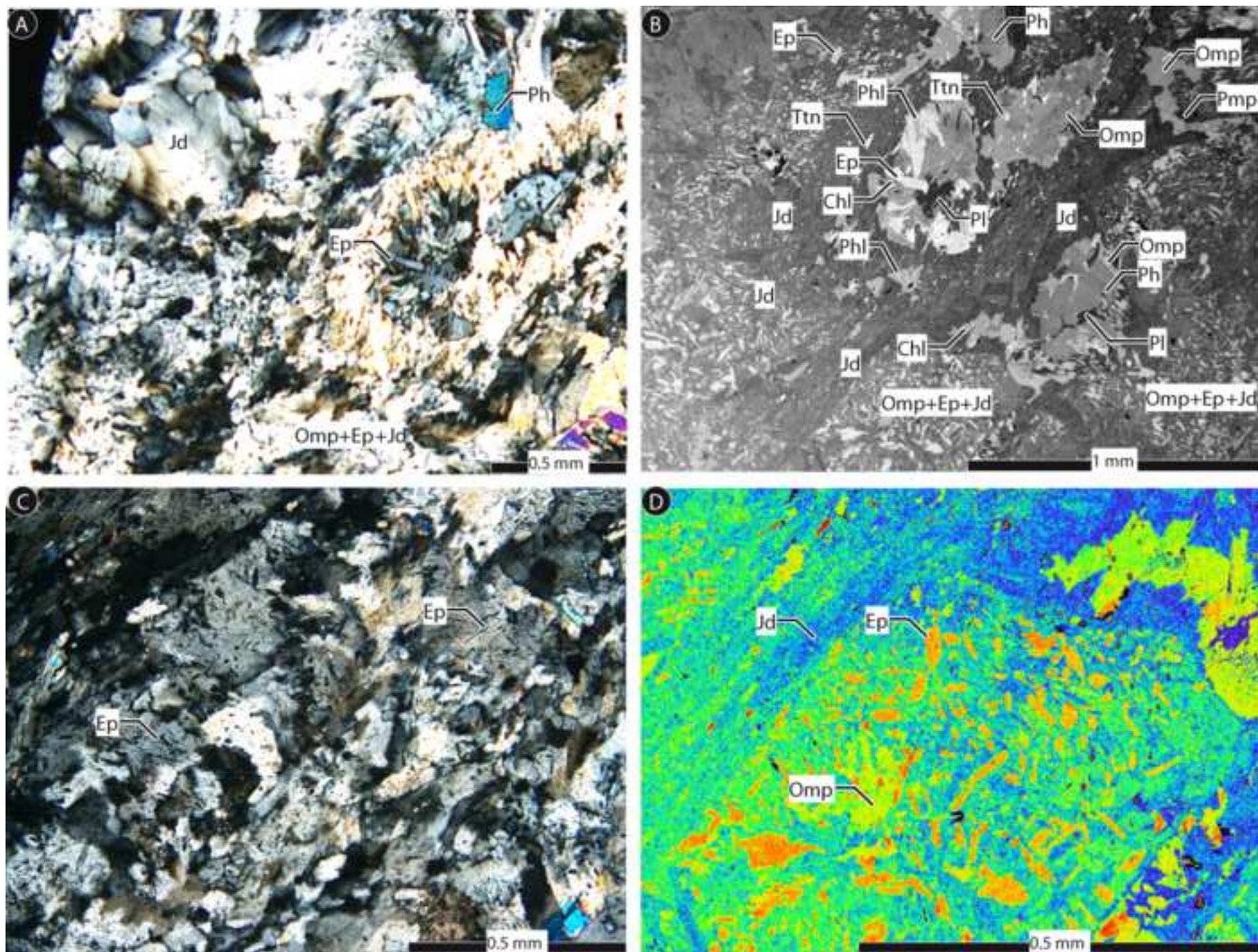


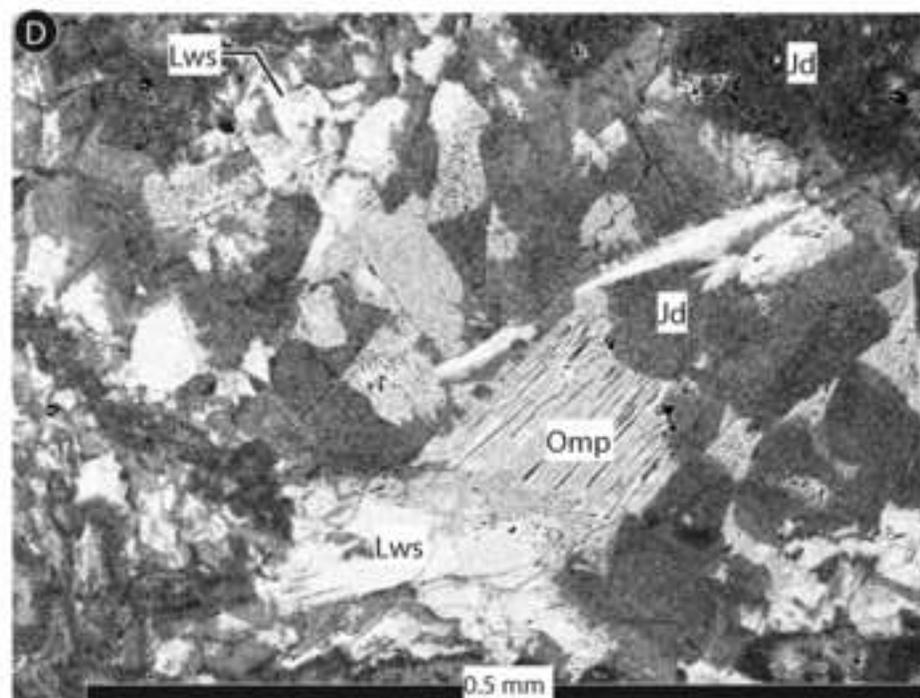
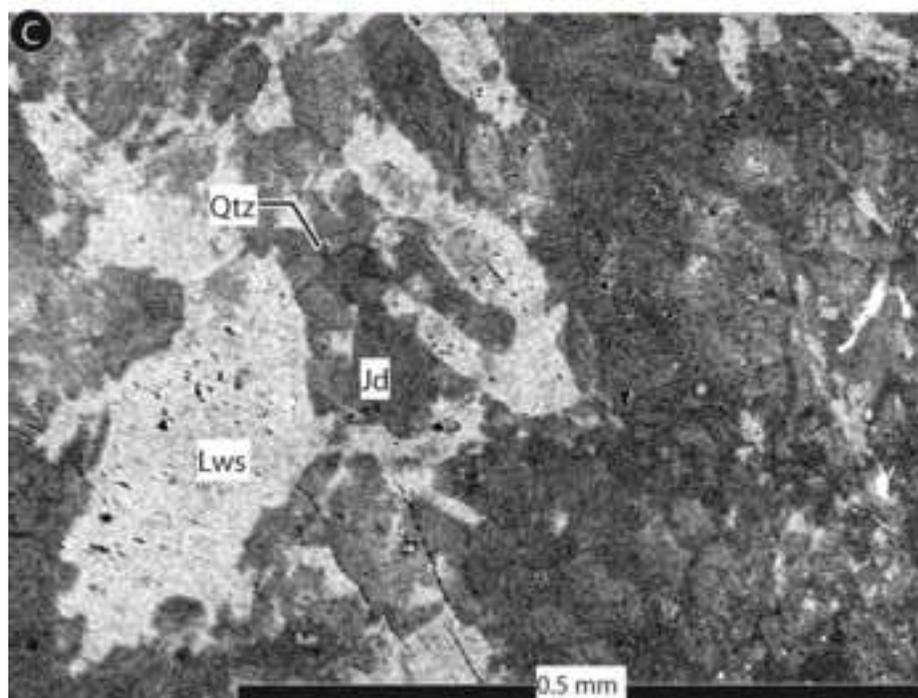
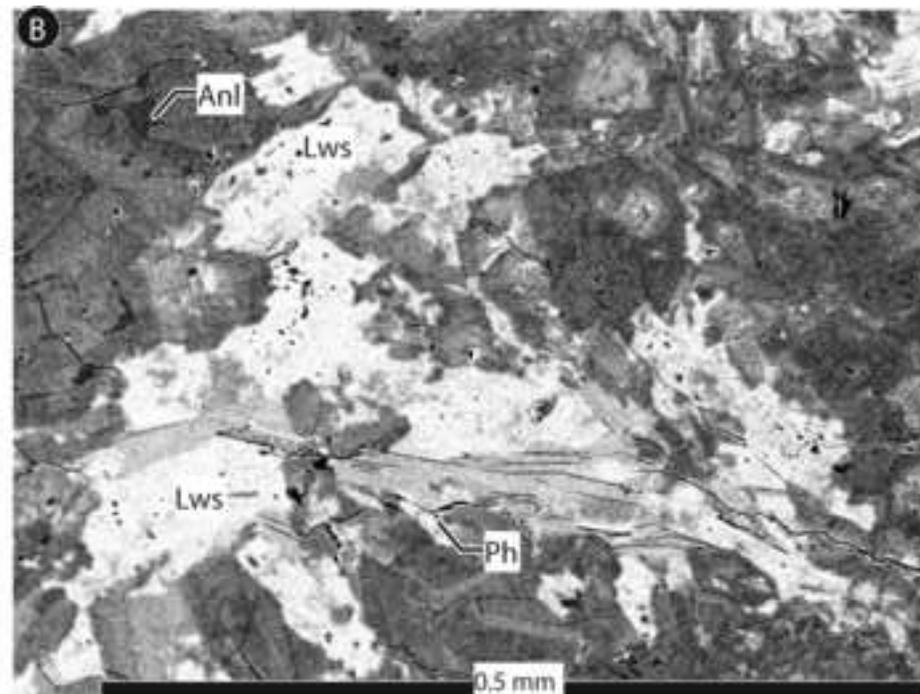
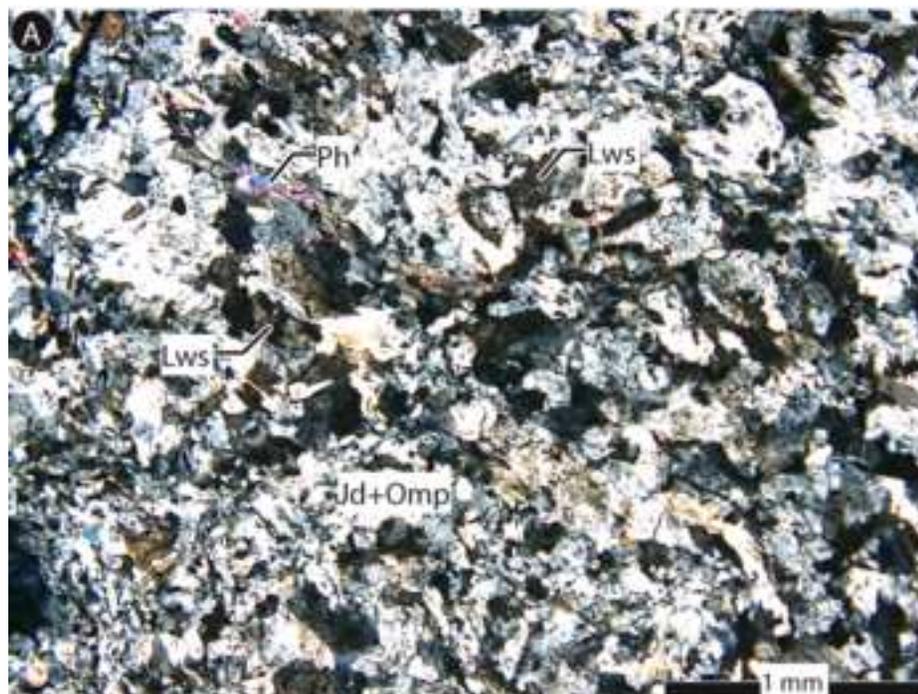
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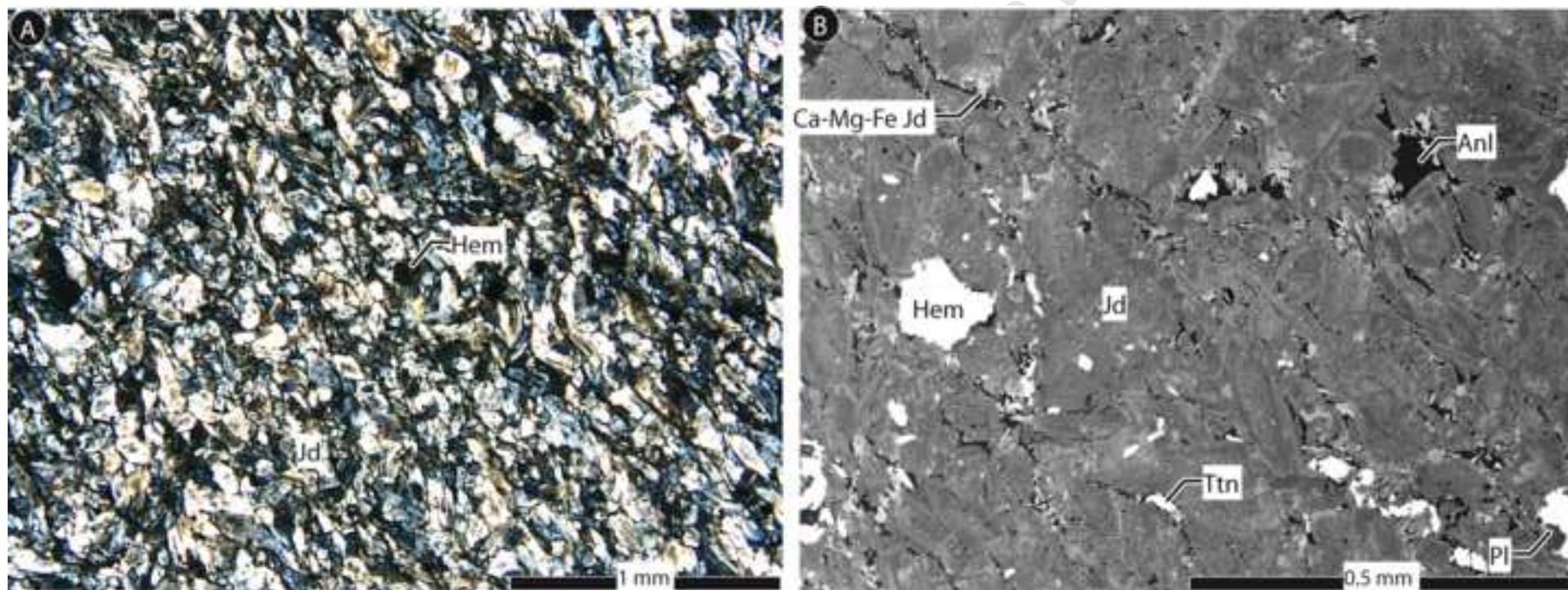


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